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**Partitioning of oblique convergence coupled to the fault locking
behavior of fold-and-thrust belts: evidence from the Qilian Shan,
northeastern Tibetan Plateau**

Mark B. Allen¹, Richard J. Walters², Shuguang Song³, Christopher Saville¹, Nicola De Paola¹,
Jonathan Ford¹, Zhenxing Hu⁴ and Wenli Sun⁴

¹ Department of Earth Sciences, University of Durham, Durham, DH1 3LE, UK

² COMET, Department of Earth Sciences, University of Durham, Durham, DH1 3LE, UK

³ MOE Key Laboratory of Orogenic Belts and Crustal Evolution, School of Earth and Space
Sciences, Peking University, Beijing 100871, China

⁴ School of Earth Sciences, Lanzhou University, Lanzhou 730000, China

Corresponding author: Mark Allen (m.b.allen@durham.ac.uk)

Key Points:

- Oblique active convergence across the Qilian Shan fold-and-thrust belt occurs by strain partitioning, similar to oblique subduction zones
- Interseismic crustal velocities are consistent with oblique motion on a single detachment thrust below 26 km depth
- Geomorphic and seismicity aspects of the Qilian Shan resemble a nascent orogenic plateau.

Abstract

Oblique plate convergence is common, but it is not clear how the obliquity is achieved by continental fold-and-thrust belts. We address this problem in the Qilian Shan, northeastern Tibetan Plateau, using fieldwork observations, geomorphic analysis and elastic dislocation modeling of published geodetic data. A thrust dips SSW from the northern range front, and underlies steeper thrusts in the interior. Cenozoic thrust-related shortening across the Qilian Shan is ~155-175 km, based on two transects. Elastic dislocation modeling indicates that horizontal strain in the interseismic period is consistent with oblique slip on a single low angle detachment thrust below ~26 km depth, dipping SSW at ~17°. We suggest this detachment is located above North China Block crust, originally underthrust during Paleozoic orogeny. Horizontal shear strain is localized directly above the up-dip limit of creep on the detachment, and is coincident with the left-lateral Haiyuan Fault. This configuration implies oblique slip on the detachment below seismogenic depths is partitioned in the shallow crust onto separate strike-slip and thrust faults. This is consistent with strain partitioning in oceanic subduction zones, but has not previously been found by dislocation models of continental interiors. The marginal, strike-slip, Altyn Tagh Fault influences thrusting within the Qilian Shan for 100-200 km from the fault, but does not control the regional structure, where Paleozoic basement faults have been reactivated. The Qilian Shan resembles the main Tibetan Plateau in nascent form: active thrusts are marginal to an interior that is developing plateau characteristics, involving low relief, and low seismicity.

1 Introduction

A key question in continental tectonics is how overall plate convergence is accommodated within plate interiors. The India-Eurasia collision has been a natural focal point for debate, because of its size, total convergence and active deformation rates [*Tapponnier and Molnar, 1977; Yin and Harrison, 2000; Tapponnier et al., 2001; England and Molnar, 2005; Royden et al., 2008*]. The Qilian Shan (Figures 1 and 2) has received particular attention because it currently deforms at rates of ~20% of the overall India-Eurasia convergence rate, and is believed to be one of the youngest areas undergoing incorporation into the Tibetan Plateau [Meyer et al., 1998; Zhang et al., 2004; Yin et al., 2008a; Duvall et al., 2013; Yuan et al., 2013; Zuza et al., 2016]. It is also an example of oblique convergence during continental collision (Figure 1) [Gaudemer et al., 1995; Zheng et al., 2013a; Daout et al., 2016], where the strain is separated (partitioned) into dip-slip and strike-slip components (Figure 2). Such zones are to be expected: ~45% of active plate boundary convergence is distinctly oblique (>22° from the normal direction to the boundary) [Woodcock and Daly, 1986]. However, it is not clear how the strain partitioning operates within individual fold-and-thrust belts. For example, what controls the location of the faults that achieve the strike-slip component of the deformation?

This paper is an integrated study of the processes of oblique convergence in the Qilian Shan, based on results from fieldwork and remote sensing data, analysis of the regional topography and seismicity, and elastic dislocation modeling of published geodetic data. In particular, we investigate how strain partitioning takes place in a region previously proposed to be underlain by a regional low-angle thrust (Figure 3) [Burchfiel et al., 1989; Tapponnier et al., 1990; Meyer et al., 1998; Yin et al., 2008b; Guillot and Replumaz, 2013], and aim to understand how the strike-slip component of such partitioning is located within the fold-and-thrust belt.

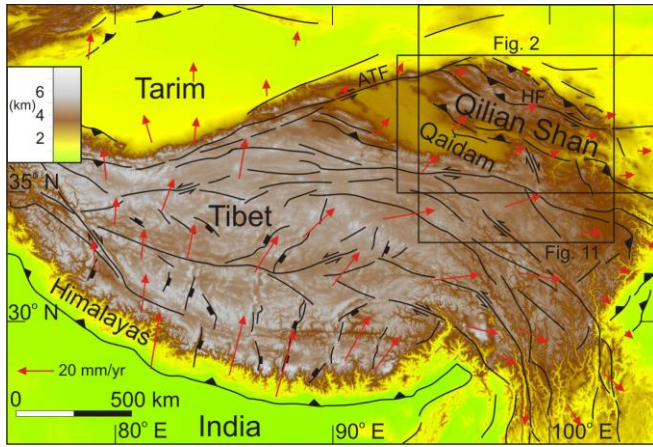


Figure 1. Summary of structures of the India-Eurasia collision, simplified from *Taylor and Yin* [2009], with GPS velocities in a stable Eurasia frame, selected from *Liang et al* [2013].

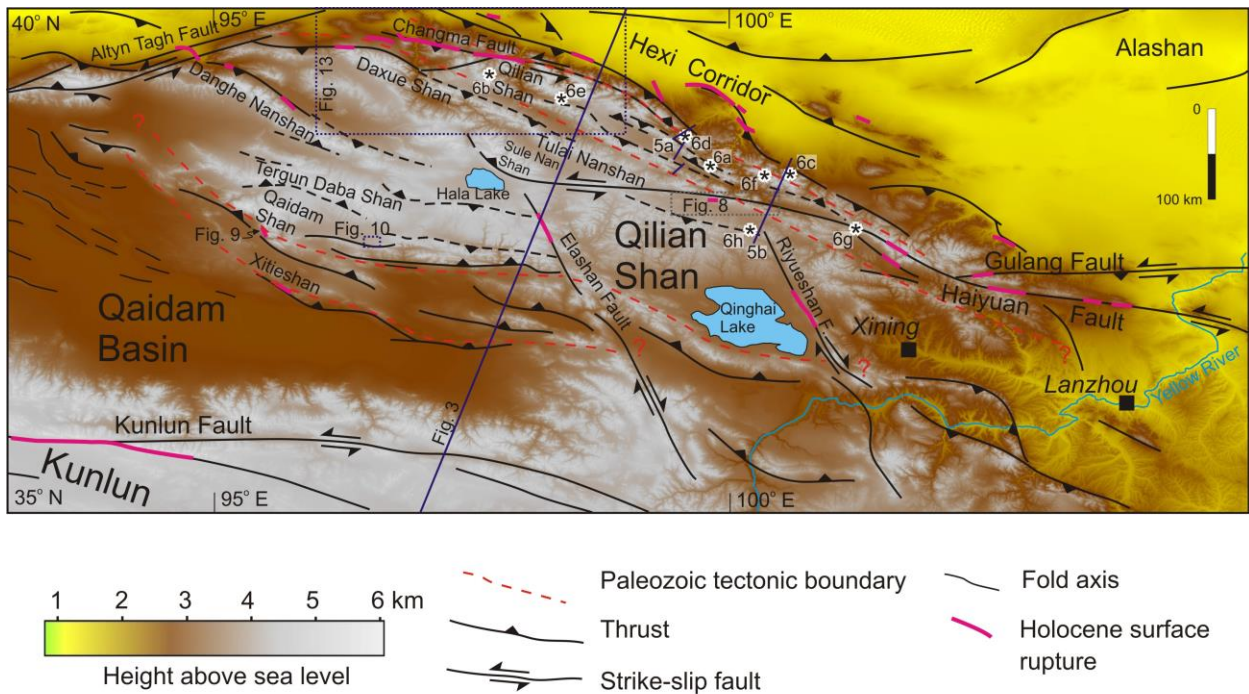


Figure 2. Major late Cenozoic structures of the Qilian Shan and Qaidam Basin, modified from *Taylor and Yin* [2009], with Hexi Corridor details from *Zheng et al* [2013b]. Earthquake surface ruptures are derived from *Tapponnier et al.*, [1990]; *Gaudemer et al.* [1995]; *Meyer et al.* [1998]; *Washburn et al.* [2001]; *Van der Woerd et al.* [2001]; *Xu et al.* [2010]; *Chen et al.* [2013] and this study.

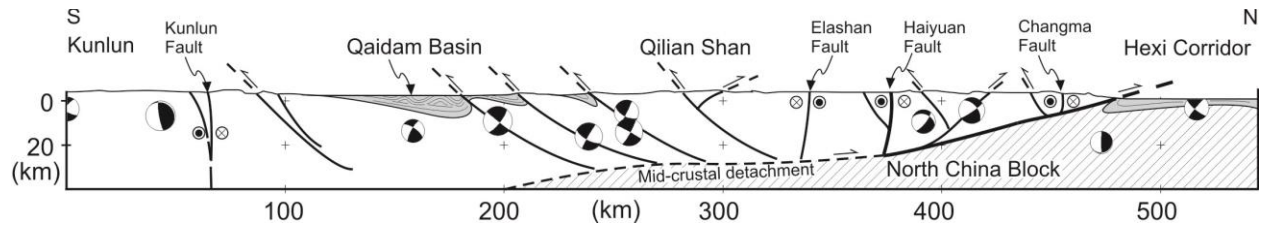


Figure 3. Interpretation of the deep structure of the Qilian Shan as detached on a regional low angle thrust, following *Burchfiel et al.* [1989] and *Meyer et al.* [1998]. Thrusts are shown as listric, and flattening onto the detachment, although this is not evident from the focal mechanisms (selected from representative data shown on Figure 4a). Bold lines indicate locked fault segments, from elastic dislocation modeling; see text for discussion. Basins shown by grey shade. Line of section shown on Figure 2.

Fieldwork in 2013 and 2014 focused on the completion of transects across the northern half of the Qilian Shan, supplemented by more reconnaissance level observations of range front geology at the southern and western sides of the region. These observations are coupled with existing 1:200,000 geological maps and satellite imagery from Google Earth to construct two regional cross-sections, at a scale of ~1:200,000, ~150 km long in total. These cross-sections help understand the location and style of deformation at present exposure levels, and help constrain the overall structure and amount of shortening. There was also focus on the evidence, or lack of it, for Quaternary and Holocene deformation in different areas.

We have used available GPS data [GEM Strain Rate Model compilation; *Kreemer et al.*, 2014 and references therein;] and simple elastic models to test whether the observed range-normal and range-parallel velocities for the Qilian Shan are consistent with interseismic strain accumulation across two shallowly-locked, localised faults; a detachment thrust and strike-slip fault respectively. This approach is a quantitative test of existing models that propose the existence of an active regional detachment thrust below the Qilian Shan [*Meyer et al.*, 1998; *Yin et al.*, 2008b; *Daout et al.*, 2016], and also allows us to investigate the geometric relationship between strike-slip and dip-slip faulting within the range.

We have analysed Shuttle Radar Topography Mission (SRTM) digital elevation data (<http://www2.jpl.nasa.gov/srtm/>; <http://srtm.csi.cgiar.org>), to gain insights from the regional topography on the present regional structure, and in particular to understand the role of the strike-slip Altyn Tagh Fault in controlling the thrust structure of the Qilian Shan; this is potentially another aspect of the oblique convergence across the range [*Meyer et al.*, 1998].

2 Regional background

2.1 Structure

The Qilian Shan (shan = mountains) consists of a series of sub-parallel mountain ranges, separated from the main part of the Tibetan Plateau by the Qaidam Basin. Each sub-range within the Qilian Shan is typically 10s of kilometres across and 100s of kilometres long (Figure 2). The name Qilian Shan is applied to the whole region between the Qaidam Basin and the discontinuous basins of the Hexi Corridor to its north, but also specifically to the northernmost of these individual ranges, and to a single peak. Present summit elevations across the Qilian Shan region are commonly >5000 m above sea level, with valley floors at ~2000-4000 m. Individual

120 ranges are usually bounded by Cenozoic thrusts [Tapponnier *et al.*, 1990; Meyer *et al.*, 1998; van
121 der Woerd *et al.*, 2001].

122 Present crustal thicknesses reach ~60-70 km in the Qilian Shan, contrasting with ~55 km
123 for the interior of the Qaidam Basin and 45-50 km for the Tarim Basin and Hexi corridor to the
124 northwest and north of the Qilian Shan [Wang *et al.*, 2013; Tian and Zhang, 2013; Feng *et al.*,
125 2014].

126 Basement of the Qilian Shan consists largely of rocks which formed in arcs and
127 accretionary complexes, as well as an ultrahigh pressure metamorphic belt. These units
128 collectively underwent a long evolution of oceanic subduction through to continental collision
129 between the Qaidam Block and the Alashan (western part of the cratonic North China Block),
130 from ~520-400 Ma [Gehrels *et al.*, 2003; Xiao *et al.*, 2009; Song *et al.*, 2013; 2014; Wu *et al.*,
131 2016]. There is also Proterozoic basement, both within the Qilian Shan, and probably underlying
132 the Qaidam Basin [Lu, 2002; Xu *et al.*, 2015]. The central part of the Qilian Shan is underlain by
133 Proterozoic basement and sedimentary cover and is known as the Qilian Block; the more
134 juvenile, accretionary crust to its north is grouped as the North Qilian Orogen. Collision between
135 the Qilian Block and the arc/accretionary material of the North Qilian Orogen took place in the
136 Early Paleozoic [Song *et al.*, 2013]. Paleozoic tectonic boundaries trend WNW-ENE, parallel to
137 the present structural and geomorphic grain of the Qilian Shan (Figure 2). Precambrian basement
138 of the North China Block (Alashan) lies to the north of the Qilian Shan, underlying Mesozoic
139 basins strung out along the Hexi Corridor [Vincent and Allen, 1999].

140 Carboniferous and Permian rocks in the Qilian Shan include both marine and non-marine
141 successions, indicating that the region was close to sea level at this time [e.g. Wang and Yu,
142 1995]. Carboniferous strata include coal measures [Geological Bureau of Qinghai Province,
143 1968a]. Mesozoic strata are typically non-marine [Vincent and Allen, 1999; Horton *et al.*, 2004],
144 including Triassic fluvial strata and Jurassic coal measures typical of other parts of northern
145 China. Cretaceous sandstones are fluvial red beds, marking more arid conditions. Cenozoic strata
146 are broadly similar to the Cretaceous rocks.

147 Cenozoic deformation and consequent exhumation may have begun as early as the
148 Eocene, i.e. not long after initial India-Eurasia collision [Clark *et al.* 2010; Zhuang *et al.*, 2011],
149 but rates accelerated and wider areas were affected by ~15 Ma or later [Palumbo *et al.*, 2009;
150 Zheng *et al.*, 2010; Lease *et al.*, 2011; Zhang *et al.*, 2012; Duvall *et al.*, 2013; Yuan *et al.*, 2013;
151 Liu *et al.*, 2017]. Such a history does not suggest a smooth outward progression to the growth of
152 the Tibetan Plateau over time, but it is broadly consistent with the northeastern part of the
153 plateau being relatively young [Tapponnier *et al.*, 2001; Yuan *et al.* 2013].

154 Late Cenozoic and active faults are shown on Figure 2, adapted from Taylor and Yin
155 [2009]. Thrusts typically lie at the foot of and parallel to the major ranges that form the Qilian
156 Shan, and trend ~WNW-ESE. There are up to seven main ranges, which have more individual
157 expression in the wider, western part of the region (Figure 2). From north to south these are the
158 Qilian Shan (*sensu stricto*), Daxue Shan-Tulai Nanashan, Sule Nanshan, Danghe Nanshan,
159 Tergun Daba Shan, Qaidam Shan and a lower elevation, discontinuous series of ridges along the
160 northern margin of the Qaidam Basin (e.g. the Xitieshan).

161 Two main left-lateral strike-slip faults form first order elements in the geology of the
162 study area (Figures 1 and 2). The Altyn Tagh Fault trends at N250° W, and forms the northern
163 margin to the Qilian Shan. The active and Holocene slip rate is ~10 mm yr⁻¹ and decreases to the
164 ENE, based on both GPS data and offsets of geomorphic features [Zhang *et al.*, 2004; Cowgill *et al.*
165 *et al.*, 2009]. Total offset has been estimated at ~350 km by several groups [e.g. Zhang *et al.*,

2001]. Thickening of the crust within the Qilian Shan has been interpreted as coupled to propagation of this fault, in a giant “horsetail” splay [Meyer *et al.*, 1998].

The Haiyuan Fault runs for ~850 km at N110° E, through the central and eastern Qilian Shan and eastwards (Figure 2), and is proposed as the main strike-slip component of overall strain partitioning [Gaudemer *et al.*, 1995]. Active slip rate estimates from GPS and InSAR are in agreement at 4-6 mm yr⁻¹, [Zhang *et al.*, 2004; Cavalie *et al.*, 2008; Jolivet *et al.*, 2012; Zheng *et al.*, 2013a]. Estimates for the total offset vary widely, from 10-95 km, based on displaced geological and geomorphic markers [Burchfiel *et al.*, 1991; Gaudemer *et al.*, 1995; Ding *et al.*, 2004]. The western end of the Haiyuan Fault occurs where it bends northwards into the Sule Nanshan range (Figure 2); this region includes the single highest peak in the Qilian Shan, Kangze’Gyai (>5800 m).

Other strike-slip faults cutting the Qilian Shan and adjacent areas to the south are the right-lateral Elashan and Riyueshan faults (Figure 2), described as having ~10 km of offset each, and likely to be moving at ~1 mm yr⁻¹ since ~10 Ma [Yuan *et al.*, 2011].

2.2 Seismicity

Several $M > 7$ earthquakes have affected the area in historic times (Figure 4a). These events include the 1927 Gulang event, which killed ~40,000 people, and, just to the east of the area covered in Figure 4a, the 1920 Haiyuan earthquake, which killed >200,000 [Gaudemer *et al.*, 1995]. These large events are concentrated on the northern side of the Qilian Shan and adjacent Hexi Corridor [Xu *et al.*, 2010]; Lee *et al.* [1976] noted that the historical earthquake record may have a recording bias towards these areas, because of their relatively high population density.

Figure 4a shows focal mechanisms for the study area, derived from a combination of body wave modelling [Molnar and Lyon-Caen, 1989; Chen *et al.*, 1999; Elliott *et al.*, 2010] and from the Global CMT catalogue (<http://www.globalcmt.org>). Thrust events are concentrated at lower elevations around the margins of the Qilian Shan, but are scarce in the range interior. Low angle events (<30° dip) are rare (Figure 4b). Clusters of earthquakes occurred in 2008 and 2009 at the southern side of the Qilian Shan (Figure 4a), which, along with other events, make this an unusually high seismicity area [Elliott *et al.*, 2011; Chen *et al.*, 2013]. Whilst most of these events are thrusts, three are strike-slip events, and are discussed in section 3.1.2. Earthquakes with $M > 4$ are also shown on Figure 4a, from the NEIC catalog (<http://earthquake.usgs.gov/earthquakes/search/>), and quoted with epicentre determinations typically of a few tenths of a degree, i.e. several 10s of km.

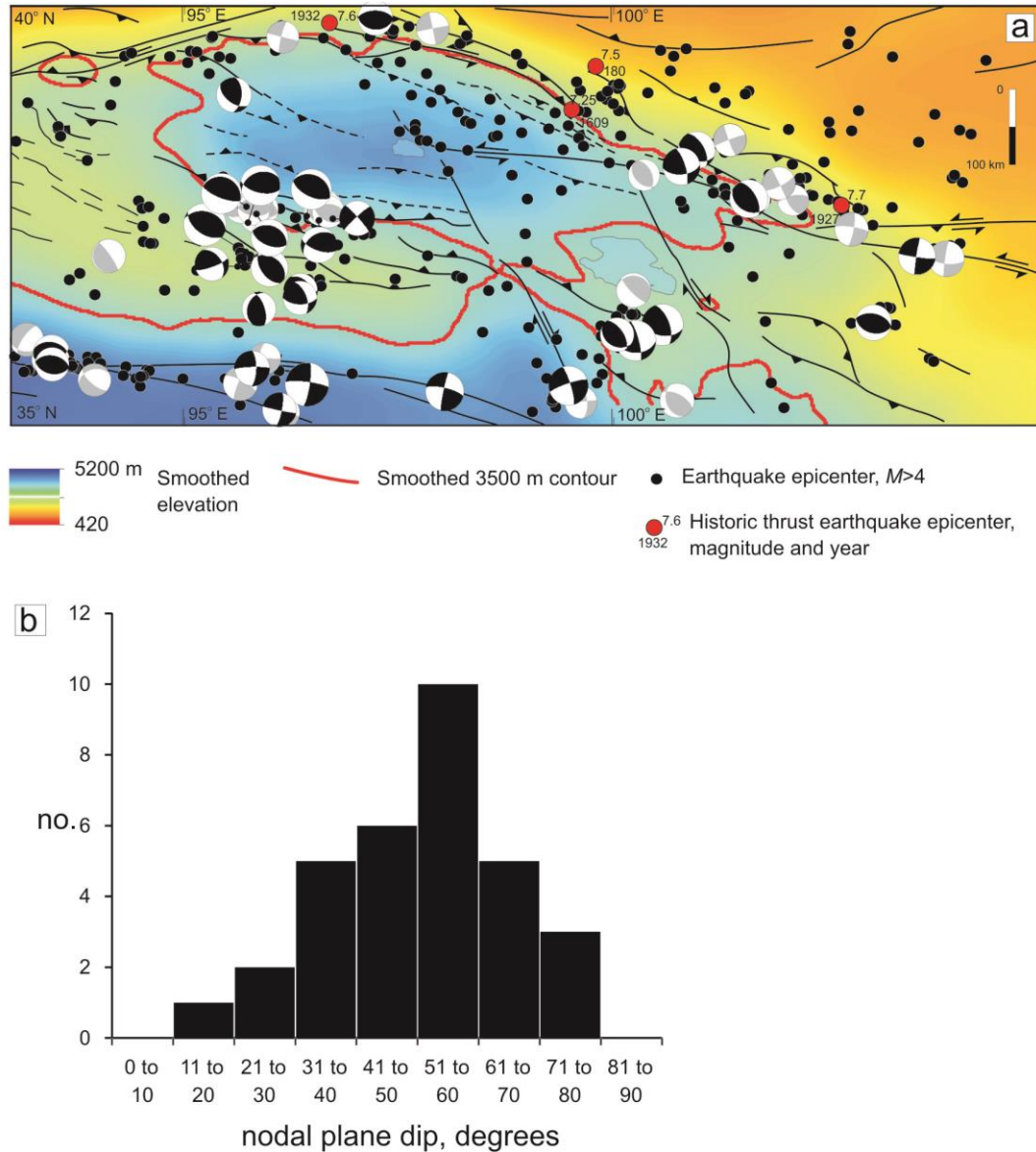


Figure 4. a) Seismicity of the Qilian Shan, Qaidam Basin and adjacent regions. Black focal mechanisms are from body wave modelling [Molnar and Lyon-Caen, 1989; Elliott et al., 2010], or the Global CMT catalog where $M > 5.3$ and there is $> 70\%$ double couple. Grey focal mechanisms are from the Global CMT catalog where $M < 5.3$ and/or there is $< 70\%$ double couple. See Supporting Information Tables 1 and 2. Epicentres of $M > 4$ events are from the NEIC catalog; historic thrust epicentres are from Xu et al. [2010]. Regional topography is smoothed over a 75 km radius moving window from SRTM data; b) histogram of thrust focal mechanism dip angles from the Qilian Shan, taken from sources listed for Figure 4a.

Several of the largest historic earthquakes produced surface ruptures which have been studied in the field [Peltzer et al., 1988; Tapponnier et al., 1990; Gaudemer et al., 1995; Meyer et al., 1998; Hetzel et al., 2004; Champagnac et al., 2010; Xu et al., 2010]. These are concentrated at the margins of the Qilian Shan, especially at the edge of the Hexi Corridor on its northern side, but also in the west, typically within ~ 100 km of the Altyn Tagh Fault (Figures 2

and 4a). Most of these events are associated with thrust faulting, but strike-slip events have been recorded, and even an oblique normal rupture associated with the 1954 Shandan earthquake in the Hexi Corridor [Xu *et al.*, 2010].

3 Observations

3.1 Structure from fieldwork and remote sensing analysis

New observations in this section add to our understanding of the distribution of strain across the Qilian Shan, and the relationship of the late Cenozoic deformation to the older structures and stratigraphy in the region.

3.1.1 Contractional deformation

Figure 5 shows two NNE-SSW cross-sections for the northern part of the Qilian Shan, centred around $\sim 99.5^\circ$ E and 100.5° E, i.e. >100 km east of sections presented in Meyer *et al.* [1998] and Şengör and Okurogullari, [1991].

Lower Paleozoic rocks are a mixture of metasediments, volcanics, granitoids, ophiolitic assemblages (e.g. gabbro and serpentinite) and high pressure metamorphic rocks [Song *et al.*, 2013]. These lithologies are typically foliated and/or intensely folded (Figure 6a), which is absent from overlying Carboniferous and younger sedimentary rocks at present exposure levels, and so constrains the Early Paleozoic age of the penetrative fabrics. Proterozoic basement and a mixed carbonate-clastic Upper Proterozoic sedimentary cover crop out in the southern part of the area covered by the section line. The valley north of these Proterozoic rocks marks the suture between the Qilian Block and the North Qilian Orogen [Song *et al.*, 2013] (Figure 5b).

Many of the regional valleys between the prominent, linear, ridges of Paleozoic/Precambrian outcrops are large scale synclines of Triassic-Cenozoic strata, overthrust from one or both margins (Figure 6b). There is a strong correspondence between the traces of major Paleozoic structural boundaries and the location of the main Cenozoic thrust faults (Figure 2) [Taylor and Yin, 2009; Song *et al.*, 2013; Wu *et al.*, 2016].

The northernmost range front represents a major thrust boundary, displacing the Qilian Shan over the Hexi Corridor to the north [Zuza *et al.*, 2016]. Thrust planes exposed within 5 km of the range front typically dip at $\leq 35^\circ$ (Figure 6c), which is a lower angle than the thrusts to the south. The mountain front is not linear, but contains salients and re-entrants up to 15 km across [Geological Bureau of Qinghai Province, 1968a] suggesting low angle thrusting.

Within the sections of Figure 5, to the south of the northern marginal thrust, there is no consistent structural vergence: thrust dip, structural relief and the degree of folding in the synclines are roughly equal on the north- and south-facing range fronts. There is no large-scale imbrication of the Triassic-Cenozoic strata, or structural windows or klippen that would indicate large-scale horizontal transport of thrust sheets on Alpine or Himalayan scales (i.e. many 10s of km).

The presence of folded and faulted Cretaceous and Tertiary strata, within the synclines and in the footwalls of the major thrusts, indicates the Cenozoic age of brittle thrusting. A dataset of these brittle thrust faults, recorded predominantly along the transects of Figure 5, has an average strike of 291° and dip of 41° (Figure 7). This average is of similar strike but lower dip than Paleozoic foliations in the same localities (295° ; 77°). Foliation data from the Xitieshan area of the southern range margin of the Qilian Shan, have a distinctly different mean orientation of $345^\circ/88^\circ$ E (Figure 7).

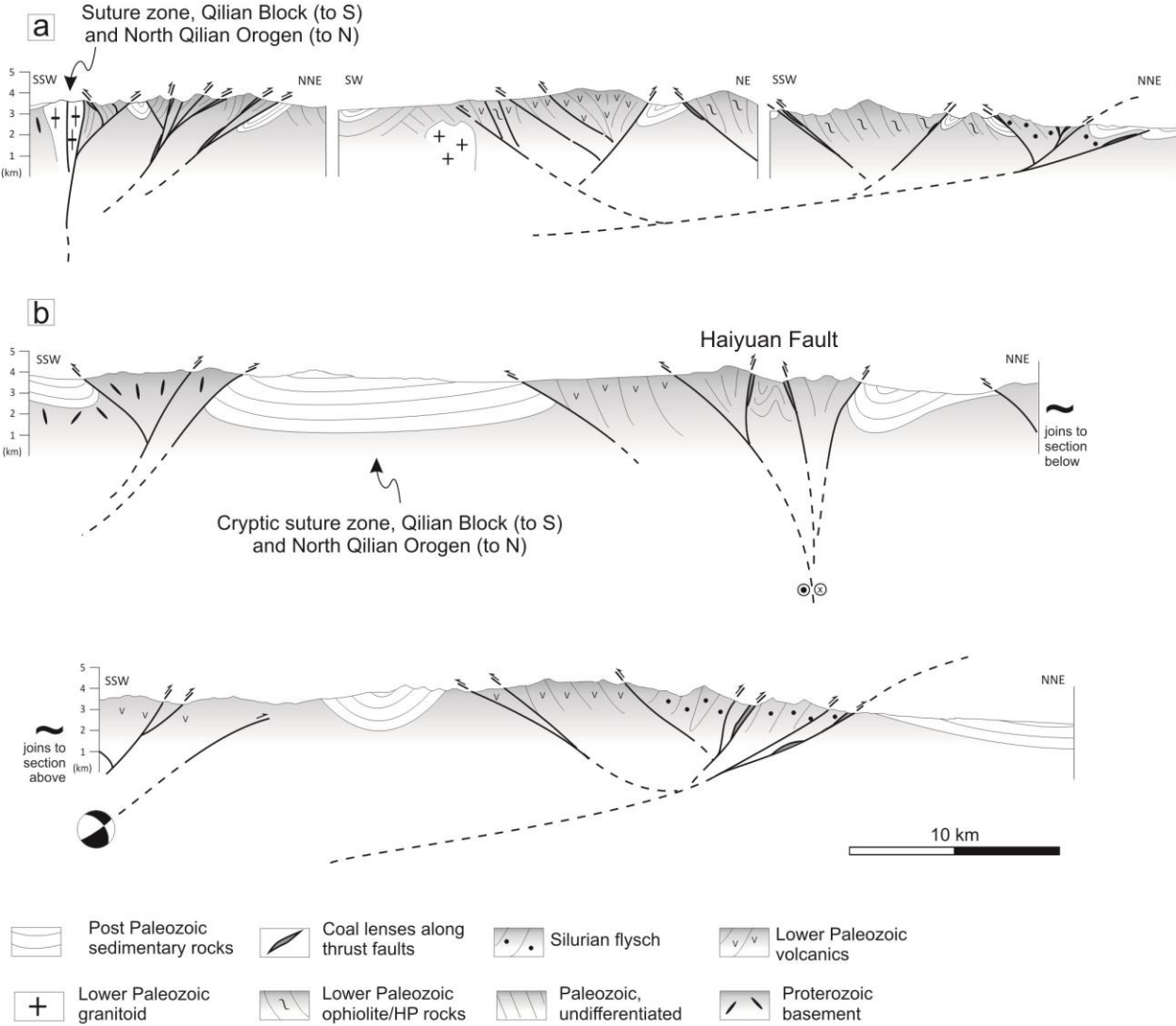
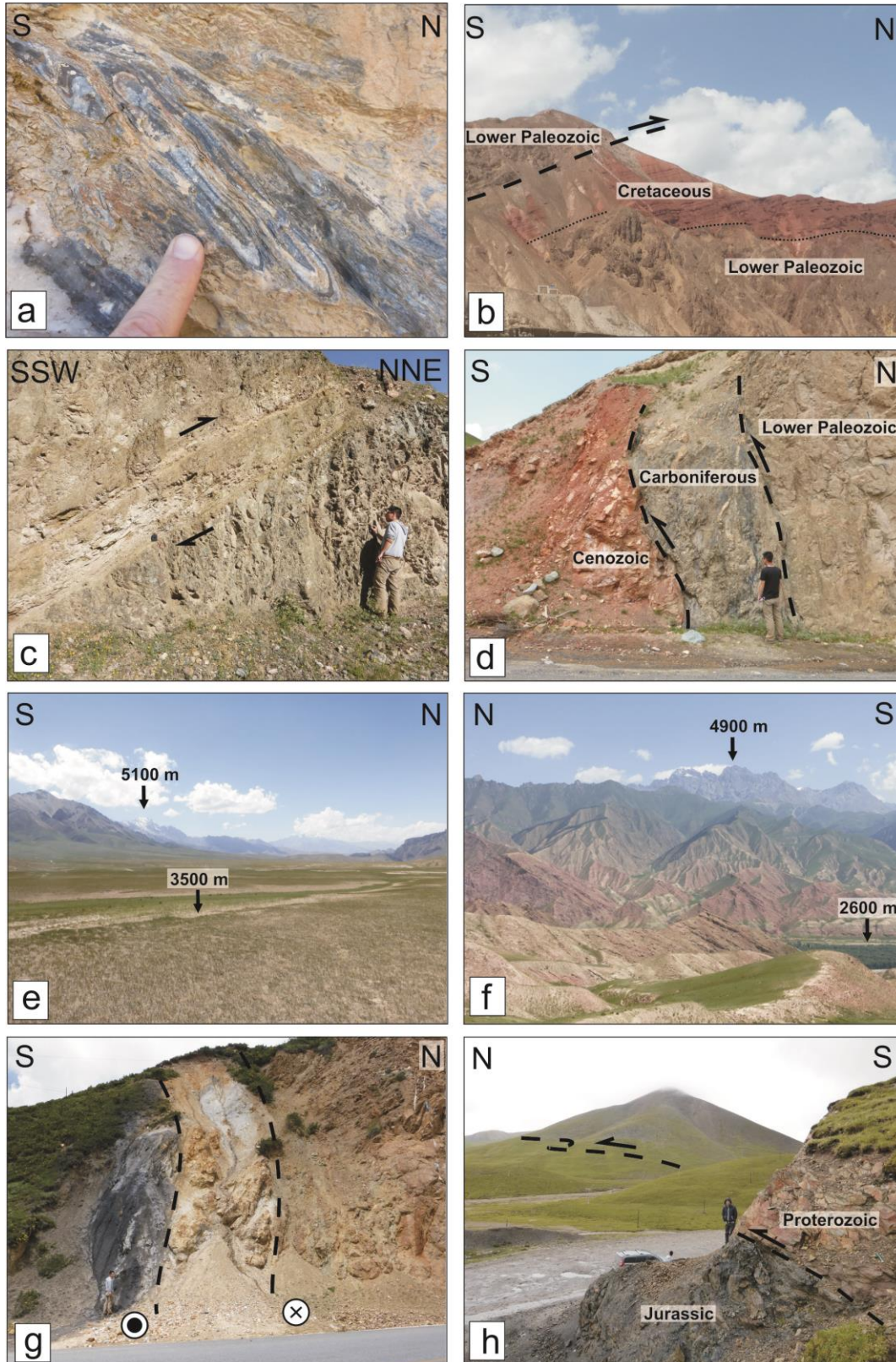


Figure 5. Cross-sections through the northern Qilian Shan, based on original fieldwork and data from *Geological Bureau of Qinghai Province* [1968a, 1968b]. Location shown on Figure 2.

Coal-bearing strata are commonly exposed in slivers along the fault planes, such that these rocks appear to lubricate many of the major thrusts (Figures 5 and 6d). These slivers are mapped as Carboniferous in age, consistent with the coal seams of that age in the region, but it is possible that some are Jurassic, given that Jurassic strata are also coal-bearing, albeit more restricted in their distribution. Fault gouge and cataclasite breccia are common, but there are not exposures of ductile shear zone lithologies such as mylonite.

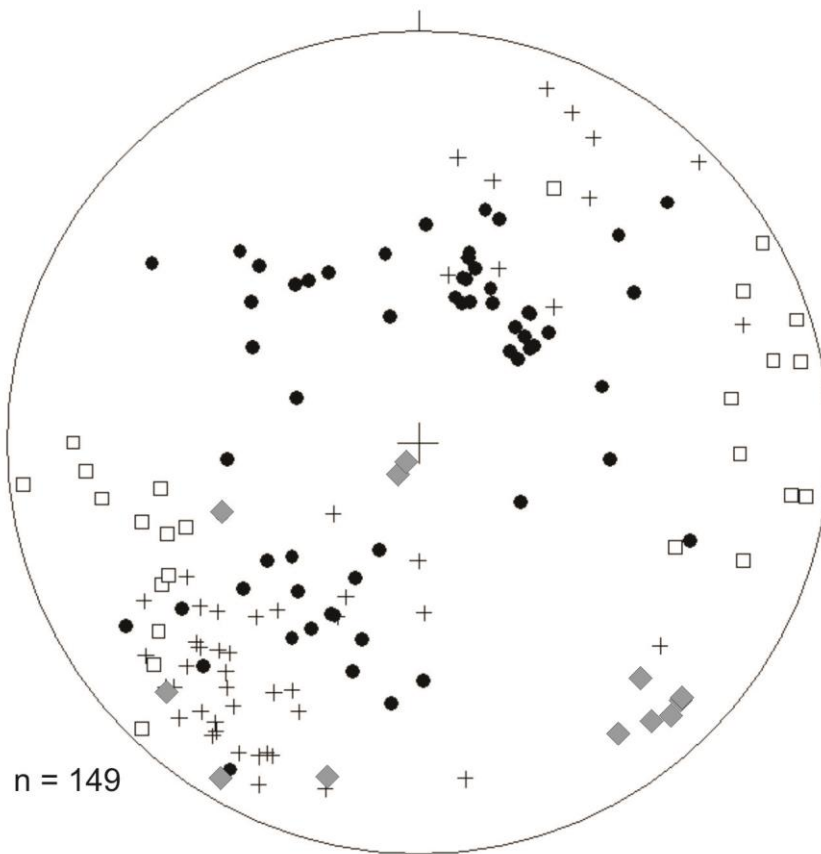
Although many range fronts are linear and sharp, there is a lack of unequivocal evidence for active thrusting along the range interiors, in contrast to the main north and south range fronts of the Qilian Shan, where active thrusting is well-documented, albeit in discontinuous segments

278 [Tapponnier *et al.*, 1990; Meyer *et al.*, 1998; Hetzel *et al.*, 2004; Yin *et al.*, 2008a; Xu *et al.*,
279 2010; Chen *et al.*, 2013].
280



281

Figure 6. Field photographs of Qilian Shan structures and landscapes, located on Figure 2: a) isoclinal folds in Lower Paleozoic blueschists; b) brittle thrust placing Paleozoic basement rocks over Cretaceous strata; c) brittle thrust with cataclasite breccia, within Ordovician volcanics at the northern margin of the Qilian Shan; d) thrusting Lower Paleozoic and Cenozoic strata, with an intervening smear of Carboniferous coal-bearing shales; e) typical landscape within the interior of the Qilian Shan, showing the relatively low relief; f) higher relief of the Qilian Shan at its northern range front; g) fault-bound slivers within the Haiyuan Fault; h) thrust at the northwestern end of the Riyueshan Fault.



Circles - poles to brittle thrust planes; Crosses - poles to Paleozoic foliations; Squares - poles to Paleozoic foliations from the UHP belt, Xitieshan; Diamonds - slickenlines from the Haiyuan Fault.

Figure 7. Stereoplot of structural data from the Qilian Shan, illustrating the variation between thrust planes of assumed Cenozoic age, and penetrative fabrics within Paleozoic rocks. Foliations in the Xitieshan region at the southern side of the Qilian Shan (Figure 2) have a more north-south orientation than elsewhere. Haiyuan Fault slickenlines are predominantly low plunge, supporting strike-slip motion, but a subset indicates dip-slip. Haiyuan Fault data are from localities shown on Figures 8 and 6b.

The majority of thrusts in the south of the Qilian Shan are directed southwards, towards the interior of the Qaidam Basin (Figure 3; see also *Yin et al.*, 2008a). Combined with the northwards thrusting at the northern margin of the range, noted above, this produces an overall divergence to the pattern of thrust transport, resembling a large scale version of the individual thrust ridges shown on Figure 5.

Triassic-Cenozoic basin fill is typically ≥ 3 km thick in the intermontane basins [e.g. *Geological Bureau of Qinghai*, 1968a]. Range crests are presently ~ 1 -2 km above the valley floors within the Qilian Shan (Figure 6e-f). These values give a rough estimate of the minimum thrust throw at each range front of ~ 5 km, using an assumption that the exposed Lower Paleozoic basement was below the level of the basal Triassic rocks at the start of Cenozoic deformation. Mesozoic deformation cannot be discounted, so that not all the throw is definitely Cenozoic in age. However, although there are local unconformities within the Mesozoic section, there is no evidence of shortening and deformation on the magnitude of the Cenozoic tectonics. Folding of Mesozoic and Cenozoic strata indicates strain within the intermontane basins, although not on the scale of the thrust ranges. Typical shortening across each of these basins in Figure 5 is on the order of 10-15%, derived by simple unfolding of the projected basal unconformity for each basin. We estimate regional shortening in the Discussion.

3.1.2. Strike-slip deformation

We examined the Haiyuan Fault between 99.5° and 100.5° E, in a region further west than previous studies of this structure. The fault in this region includes an eastern segment which trends sub-parallel to the Paleozoic basement grain, passing westward into a segment at $\sim 100.2^\circ$ E, which cuts obliquely across the basement and folded Mesozoic rocks (Figure 8a). The fault zone contains anastomosing lenses of contrasting lithologies, at varying scales upwards from meter-sized lozenges (Figure 6g) to km-scale slivers. Weak lithologies such as shale, coal and serpentinite are common, deformed by intense brittle fault networks. Sub-horizontal slickenlines are common within the Haiyuan Fault, but rare along other major structures shown within Figure 5. Even within the Haiyuan Fault there are dip-slip slickenlines, indicating thrust motion (Figure 7). We cannot discount that at least some of these dip-slip structures relate to pre-Cenozoic deformation, but as noted, Paleozoic rocks are typically deformed with penetrative fabrics not seen in the post-Paleozoic section. The width of the entire Haiyuan Fault zone is ~ 10 km in this region, based on the presence of sub-vertical faults with sub-horizontal slip across this distance.

Offset bedrock geology can be matched across the Haiyuan Fault in the region of 100° E, i.e. where the fault cuts across the Paleozoic grain, and further west than the locations of most of the previous offset estimates [*Burchfiel et al.*, 1991; *Gaudemer et al.*, 1995; *Ding et al.*, 2004]. Paleozoic and Mesozoic rock units and the faults that separate them are offset for ~ 10 km (Figure 8a), based on 1:200,000 survey mapping [*Geological Bureau of Qinghai Province*, 1968a; 1968b]. We confirmed the left-lateral slip sense and vertical orientation of this fault where it juxtaposes ophiolitic *mélange* against Permian strata (Figure 8a). This offset is distinctly lower than most previous estimates for the Haiyuan Fault (which are up to 95 km, *Gaudemer et al.*, 1995), based on a variety of approaches from bedrock offset, to the apparent left-lateral offset of the Yellow River (Figure 2). The difference between our estimate and most previous studies suggests that Figure 8a shows only part of the total offset, with major but unquantified slip occurring on other fault strands to the north and south, examined during our fieldwork (Figure 8a). An alternative but complementary explanation for our relatively low offset estimate

is that there is an along-strike gradient in slip, decreasing westwards towards the end of the Haiyuan Fault in the Sule Nan Shan (Figure 2).

Offset is harder to determine where the Haiyuan Fault is parallel to the regional bedrock structure. On the regional scale, the Haiyuan Fault also crosses the Paleozoic structural grain of the Qilian Shan at a low angle of obliquity, and from east to west passes into the interior of the range from its northern margin (Figure 2). The segments parallel to the basement fabric are effectively right-stepping restraining bends, and have been associated with thrust earthquakes (Figure 4a). This is a possible explanation for the more steeply-plunging, dip-slip slickenlines within the broad fault zone (Figure 7).

Figure 8b shows a previously unreported trace of a segment of the Haiyuan Fault, where it cuts the Quaternary sediments of an intermontane basin for a length of ~12 km. The segment trends east-west, and continues from the western end of the strand described above at 100° E. Individual stream gullies are offset left-laterally by ~35m (Figure 8c), suggesting that this offset represents cumulative Quaternary displacement over an unknown number of earthquakes, rather than a single rupture.

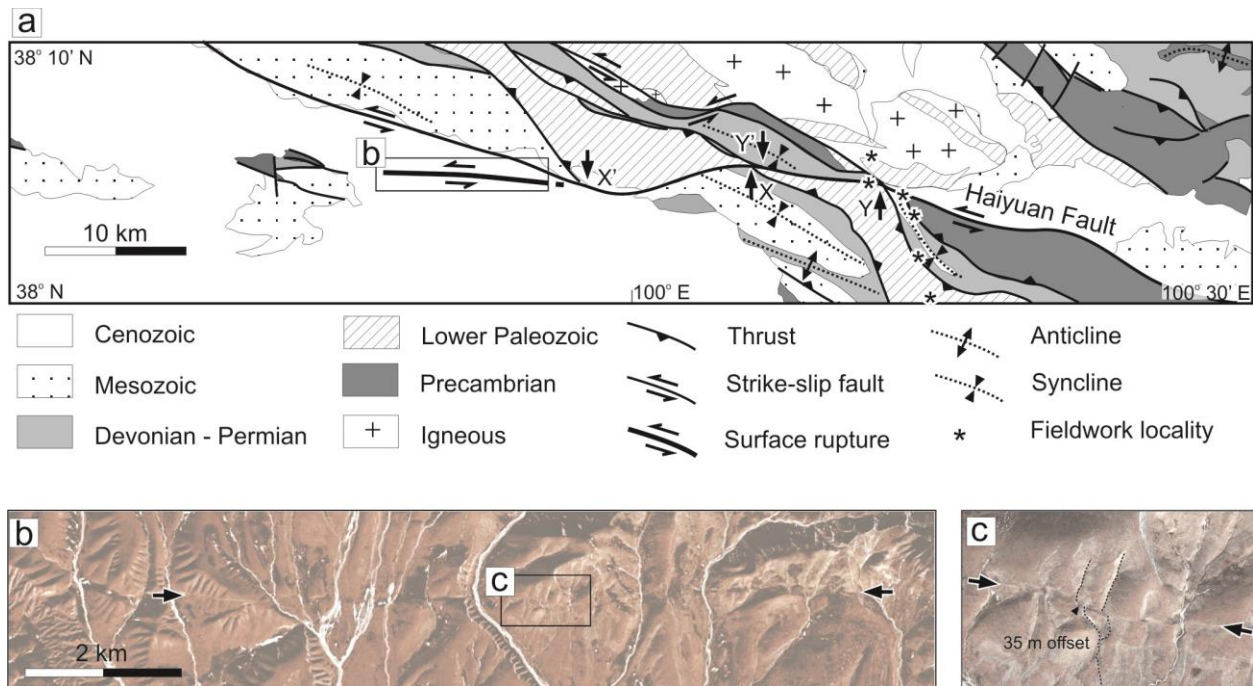


Figure 8. a) Structure of the Haiyuan Fault at ~100°E [Geological Bureau of Qinghai Province, 1968a, 1968b]; b) surface rupture across Quaternary sediments, located in a); c) close-up of ruptures in b). Quickbird satellite imagery from Google Earth (©2014 Google, ©2014 DigitalGlobe).

We report another previously-unidentified surface rupture at the southern margin of the Qilian Shan, near the Xitieshan Fault, on the Qaidam Thrust [Yin *et al.*, 2008b] (Figure 9). Here, there is a linear, sharp scarp in the surface of the alluvial fans. The rupture can be traced for ~8 km across the piedmont (Figure 9a,b), consistent with a $M \sim 6$ Holocene earthquake in this area [Wells and Coppersmith, 1994]. The slip sense is not clear and needs further work (access was not possible during our fieldwork). The regional elevation of the adjacent range to the northeast

argues for a thrust component. A strike-slip or oblique sense is possible; faint rupture patterns in the imagery suggest a right-lateral component (Figure 9c), which would be consistent to the right-lateral slip observed to the east on the Elashan and Riyueshan faults (Figure 2).

A fault trending northwest-southeast, east of Xitieshan, is visible in satellite imagery (Figure 10), but does not appear to have been mapped previously [*Geological Bureau of Qinghai Province*, 1980]. Offset sense is right-lateral, and the fault is parallel to the much larger and more prominent Elashan and Riyueshan faults to the east (Figure 2). The faulting east of Xitieshan may represent a less developed equivalent of the Elashan and Riyueshan faults, which were interpreted by *Duvall and Clark* [2010], *Yuan et al* [2011] and *Zuza and Yin* [2016] to help accommodate shortening by counter-clockwise rotation. Three strike-slip earthquake focal mechanisms corroborate the fault slip sense and show it to be an active structure (Figure 10e).

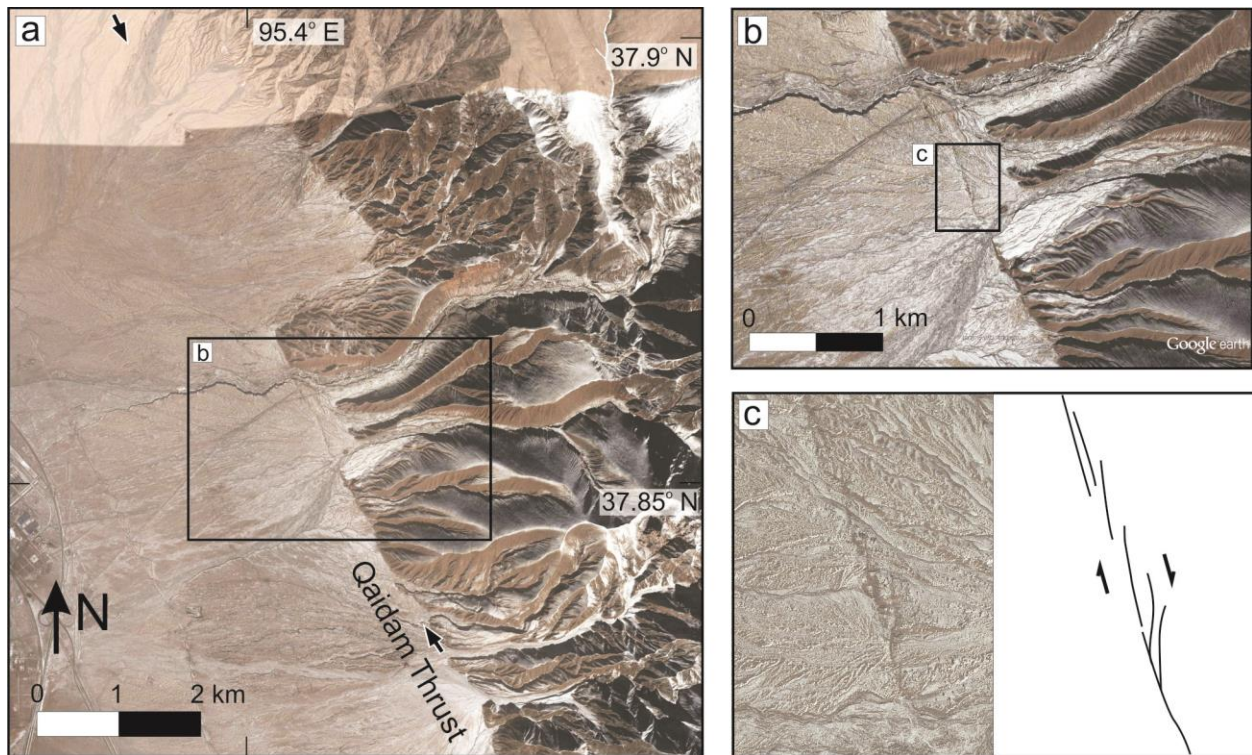


Figure 9. Quickbird satellite imagery from Google Earth (©2014 Google, ©2014 DigitalGlobe) of the Chaidanzhen fault break on the Qaidam Thrust. Location shown on Figure 2.

Yuan et al [2011] suggested that the northwestern end of the strike-slip Riyueshan Fault terminated in a thrust, but did not examine the structure in this region. We confirm that a south-dipping thrust places Proterozoic rocks over Jurassic coal measures at the end of the Riyueshan Fault (Figures 2 and 6h), but the topography is subdued, with no strong indications of Holocene activity.

In summary, left-lateral faulting within the Qilian Shan is focused on the Haiyuan Fault, as far west as its termination in the Sule Nanshan, but the fault zone is up to ~10 km wide and complex, and includes segments that both cross and lie parallel to the regional Paleozoic basement grain. Shorter right-lateral faults are also present within the Qilian Shan, additional to the Elashan and Riyueshan faults previously documented.

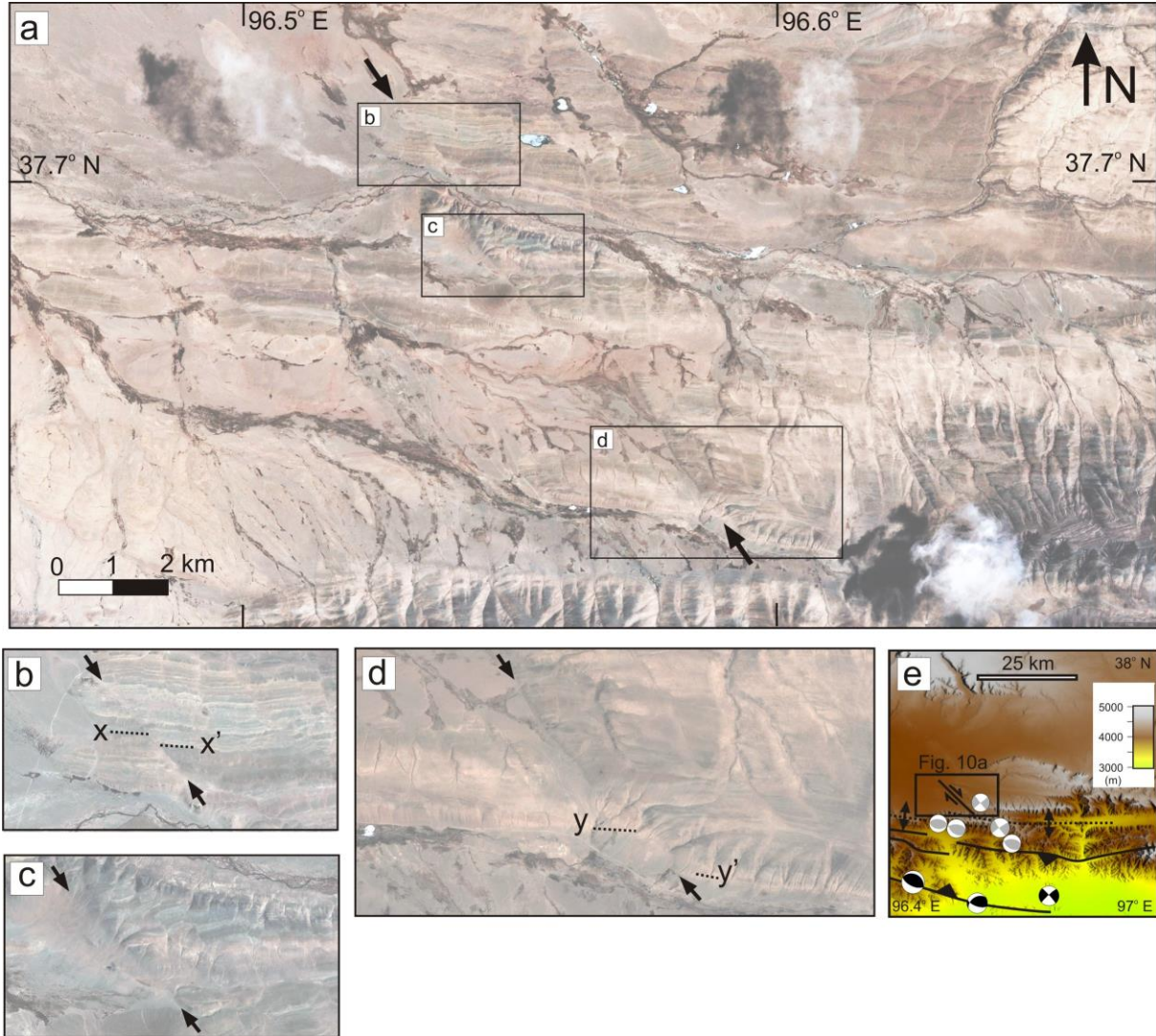


Figure 10. a-d) Quickbird satellite imagery from Google Earth (©2014 Google, ©2014 DigitalGlobe) showing right-lateral strike-slip faulting east of Xitieshan. Location shown on Figure 2; e) location of Figure 10a, and local focal mechanisms as in Figure 4a.

3.2. Constraints from geodetic data and elastic dislocation modeling

GPS studies have previously established a horizontal velocity field for our area of interest [Gan *et al.*, 2007; Liang *et al.* 2013] (Figure 11). In a Eurasian reference frame, interseismic velocities decrease from $\sim 20 \text{ mm yr}^{-1}$ in the northern Kunlun, to $\sim 2\text{-}3 \text{ mm yr}^{-1}$ in the Alashan, on the edge of the Hexi Corridor, and exhibit partitioning of oblique convergence into components that are orthogonal and parallel to regional fault trends [Zhang *et al.*, 2004; Zheng *et al.*, 2013a]. Previous studies have used elastic dislocation models to interpret the GPS data, but have mainly focused on modeling the range-parallel strike-slip component, thought to be largely accommodated on the left-lateral Haiyuan fault, and in contrast have suggested the convergent strain is broadly distributed across the Qilian Shan [Zheng *et al.*, 2013a; Zhang *et al.*, 2013]. In

these models, the investigation of model fault locations and geometries is often precluded, as these details are fixed a-priori.

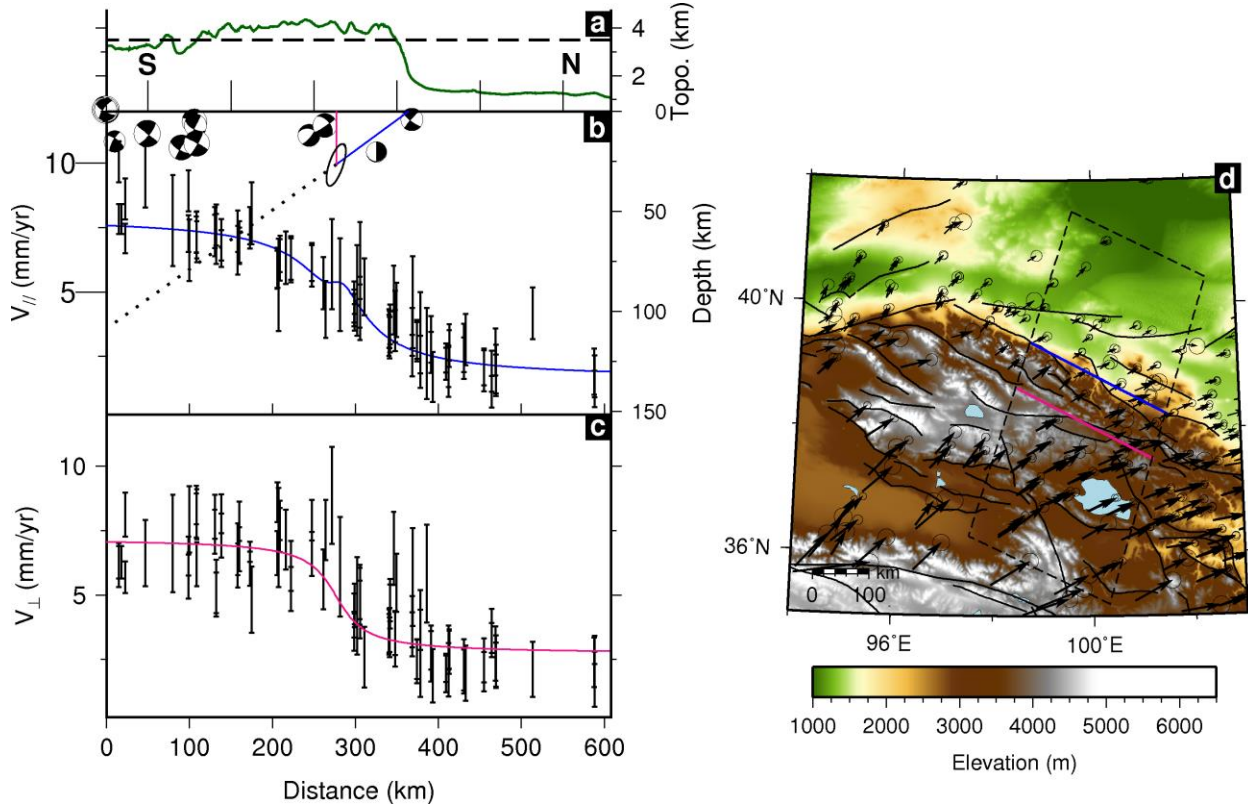


Figure 11. One-fault elastic dislocation model for interseismic deformation in the Qilian Shan. a) topographic profile from dashed box within (d), from SRTM data; b) best-fit model fault geometry, earthquake focal mechanisms and fit to data for profile-parallel velocities. Dashed black straight line shows the geometry of the single creeping detachment fault with oblique slip, and the black ellipse shows the 95% uncertainty on the location of the locked-creeping transition. Above this point, the blue and pink solid lines show the inferred geometry of the locked thrust and strike-slip faults, onto which the oblique slip is inferred to be partitioned. Focal mechanisms show location of earthquakes from Figure 4a for the region of (d). Black vertical bars show GPS-derived velocity components parallel to the section line with 2-sigma uncertainties, and blue line shows predicted velocities for the model geometry shown in the same panel.; c) Black vertical bars show velocity components normal to the section line with 2-sigma uncertainties, and pink line shows predicted velocities for the model geometry shown in (b); d) map of GPS data from Kreemer et al., [2014] compilation. Dashed box shows region of GPS data projected onto central profile line for (b) and (c) and pink and blue lines show surface projection of the strike-slip and thrust model faults.

Here we adopt a different approach, to test whether the observed range-normal and range-parallel velocities across the Qilian Shan are consistent with interseismic strain accumulation across a single, localised detachment thrust and a sub-vertical strike-slip fault respectively. In this approach, the geometry and location of the faults are free parameters, which enables us to investigate if such a simple two fault model can fit the available data, and if it can, the geometry

and location of these structures, along with the relationship between them. In particular, this is a quantitative test of existing models that propose the existence of an active regional detachment thrust below the entire Qilian Shan that accommodates the majority of the convergence across the mountain belt [Meyer *et al.*, 1998; Yin *et al.*, 2008b], as opposed to suggestions that convergence is broadly distributed across the mountain belt with little evidence of localisation [Zheng *et al.*, 2013a; Zhang *et al.*, 2013].

3.2.1 Two-fault model

First we draw a profile orientated N20E, orthogonal to regional fault trends (Figure 11a), and project published GPS velocities onto the profile line from within ~250 km, separating the profile-parallel (shortening) and profile-perpendicular (strike-slip) velocities (Figure 11b,c). We have used published GPS data from the GEM Strain Rate Model compilation [Liang *et al.*, 2013; Kreemer *et al.*, 2014 and references therein] (Figure 11d). Initially, we model the horizontal velocity components separately with two model faults, in each case using a back-slip calculation for a dislocation in an elastic half space, where the modeled fault is locked by friction above a ‘locking’ depth, and slips freely below this depth [Savage and Burford, 1973; Savage, 1983]. The model parameters for the strike-slip fault are slip rate, locking depth and location of the fault, as well as a static offset in velocity to account for a non-zero reference. For the thrust fault, we fix the surface projection of the fault to the northern range front of the Qilian Shan, and solve for slip-rate, locking depth and dip of the fault, along with a static offset in velocity to again account for a non-zero reference. The horizontal location of the locked-creeping transition is also a parameter of interest, but is not independent; it is constrained solely by the dip and locking depth. For both models, we use a weighted non-linear least squares inversion with multiple restarts to avoid local minima. We estimate uncertainties on each model using a Monte Carlo approach, perturbing the GPS data 100 times with noise based on the formal velocity uncertainties and inverting these perturbed datasets using the same process detailed above. The spread of the 100 retrieved values for each model parameter define the uncertainty on that parameter. More details of the code used are available from the authors on request.

We make one additional modification to our two-fault model: the best-fit detachment fault has a locking depth of 35 ± 15 km (2-sigma), which is much greater than the estimated 20 km seismogenic thickness [Sloan *et al.*, 2011] and would lie within the lower crust for this region. We therefore fix the depth to 20 km, with very little degradation to the fit to the data (see Supporting Information: Figures S1 & S2 and Table S3).

Our two-fault model is presented in Table 1 and Figure S1 and fits the GPS data well, with residuals of similar magnitude to the formal uncertainty on the data (~1 mm/yr). The most important results of this model are; 1) that the range-normal component of the GPS data can be explained by slip on a localized detachment at depth and; 2) that the data strongly constrain the location of the locked-creeping transition of the detachment to be coincident (within uncertainty) with the locked-creeping transition for the strike-slip fault and with the mapped trace of the Haiyuan fault. Essentially, the locked portion of the strike-slip fault lies above the locked-creeping transition of the detachment thrust (Figure 3, Figure S1b). We note that the dip angle and dip direction of the detachment thrust are not well-constrained, but although models with different dip values provide similar fit to the data (Supporting Information Figures S2 and S3), the observation of the intersection at depth of the strike-slip and detachment thrust is robust

irrespective of dip direction, for locking depths <20 km. Here we present the south-dipping solution as this seems more realistic, given the location of the Qilian Shan at the northeastern side of the Tibetan Plateau. In the *Savage and Burford* [1973] elastic model for strike-slip faulting used here, fault-parallel surface velocities do not depend at all on the dip of the strike-slip fault below the locking-depth. Therefore, the simplest interpretation of the coincidence between the locked-creeping transitions of the two model faults is that the dip of the strike-slip fault at depth is the same as the detachment, i.e. that oblique slip is localized on a single dipping detachment at depth, and is partitioned only in the seismogenic crust.

3.2.2 One-fault model

We test this interpretation formally by running one final inversion. We use the same method described above but this time constrain both strike-slip and thrust components of slip to take place on a single dipping detachment below one locking depth, simultaneously fitting both profile-perpendicular and profile-parallel GPS velocities. The joint inversion of both components of velocity means the locking depth is better constrained than in the previous inversions, so it is not fixed in this model. When compared to the two-fault model, the resulting one-fault oblique-slip model fits the data with no significant degradation to misfit, and is presented in Table 1 and Figure 11.

The best-fit oblique detachment slips below a locking depth of 26 ± 8 km, and has a dip of 17 ± 4 degrees to the SSW. This locking depth is in close agreement with estimates of the seismogenic thickness from robust earthquake depths (~ 20 km; *Sloan et al.*, [2011]), and with previous interseismic studies of the Haiyuan fault west of 104° E, [*Jolivet et al.*, 2012], which is fully locked, in contrast to the creeping section further east [e.g. *Jolivet et al.*, 2015; *Daout et al.*, 2016]. The best-fit horizontal location of the locked-creeping transition on the fault is constrained to within 9 km, and is in agreement with the mapped surface trace of the Haiyuan Fault, implying a vertical or sub-vertical structure in the seismogenic upper crust. The rate of strike-slip motion on the detachment is 4.2 ± 0.4 mm yr $^{-1}$, consistent with previous estimates for the slip-rate on the Haiyuan fault in this region [e.g. *Zhang et al.*, 2004; *Cavalié et al.*, 2008; *Jolivet et al.*, 2012; *Zheng et al.*, 2013a]. The rate of reverse dip-slip motion on the detachment is 6.7 ± 1.0 mm yr $^{-1}$. It is important to note that although we draw the shallow, locked thrust fault that accommodates this strain in the seismogenic crust as a low angle continuation of the detachment (Figure 3), which matches geologic data (Figure 5; *Zuza et al.*, 2016), this is not a requirement of the dislocation model itself. The model does not predict the structure within the locked region of the crust, i.e. whether the dislocation continues up-dip at a low angle or into one or more steeper structures. From the ratio of strike-slip and dip-slip rates (Table 1), we estimate the rake on the detachment at depth to be $\sim 58^\circ$.

The geometric relation proposed here between major strike-slip and thrust faults has been inferred from geodetic data above subduction zones with oblique convergence, and is suggested from numerical models to result from localisation of shear strain in the crust above the locked-creeping transition on the detachment [*McCaffrey et al.*, 2000]. This relationship has not been previously identified for the Qilian Shan (although it was predicted by the models of *Bowman et al.*, 2003), and is highly likely to occur in other continental fold-and-thrust belts where strain partitioning has also been recognised [e.g. *Murphy et al.*, 2014; *Silver et al.*, 2015].

Model fault type		Slip-rate (mm yr ⁻¹)	Locking Depth (km)	Dip (degrees)	Location of locked-creeping transition along profile (km)	Offset in velocity (mm yr ⁻¹)	RMS (mm yr ⁻¹)
Two faults	Dip-slip	6.0 ± 0.4	20*	13 ± 2 S **	49 ± 12**	4.9 ± 0.2	0.78
	Strike-slip	4.1 ± 0.8	18 ± 21	90*	63 ± 6	4.3 ± 0.2	1.03
One fault, two slip components	Dip-slip	6.7 ± 1.0	26 ± 8	17 ± 4 S	49 ± 9	4.8 ± 0.2	0.74
	Strike-slip	4.2 ± 0.4				4.4 ± 0.2	1.06

Table 1. Elastic dislocation model parameters for initial model with two separate faults, and for final model with one fault and oblique slip at depth. Uncertainties are at 2-sigma level. *fixed in inversion. **when depth is fixed, dip and location of fault co-vary.

3.3 Regional scale geomorphology and drainage patterns

We have examined regional patterns of topography and drainage, to add at least first order insights to our fieldwork, remote sensing and dislocation models. Smoothed elevation values (Figure 4a) and slope values (Figure 12) highlight that the highest elevation and lowest relief part of the Qilian Shan occurs in the middle of the region, centred on Hala Lake. Regional elevations decrease away from this area in all directions, including towards the Altyn Tagh Fault to the west and north.

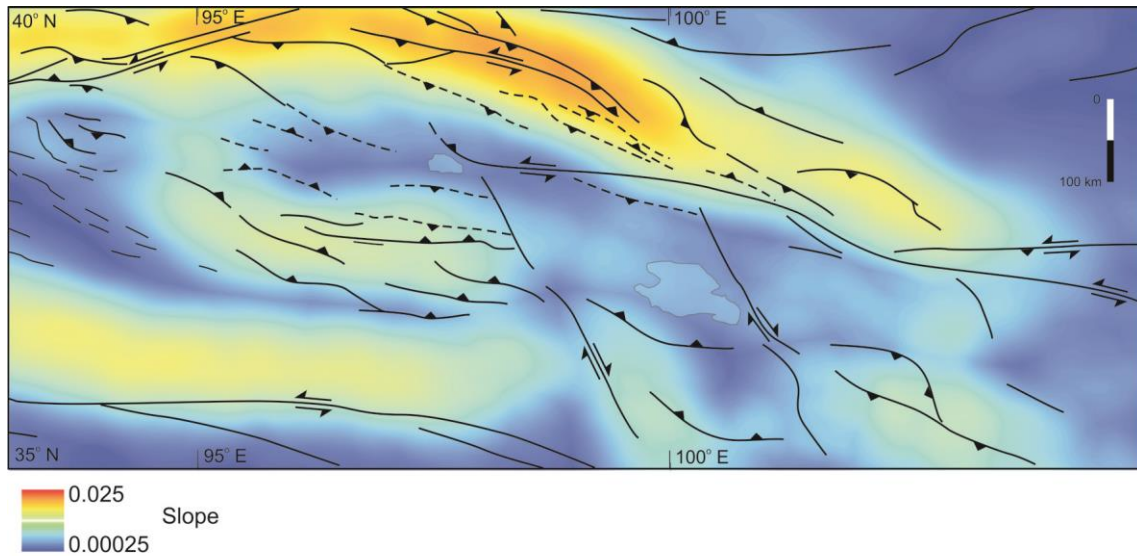


Figure 12. Regional slope, derived from a 35 km radius moving window on mean elevation smoothed over a 75 km radius moving window.

Rivers in the east of the Qilian Shan that do not drain into the Yellow River typically flow into Hala or Qinghai lakes (Figure 2). Rivers in the west flow to the WNW along the major

valleys parallel to the range fronts (Figure 13), but three of the major ones are diverted northwards before they reach the Altyn Tagh Fault. This northwards diversion produce a nested pattern to the drainage (Figure 13), with the southernmost river, the Sule He, reaching furthest west, while rivers to the east, the Beida He and Hong Shuiba He, have the shorter courses and more easterly diversion points.

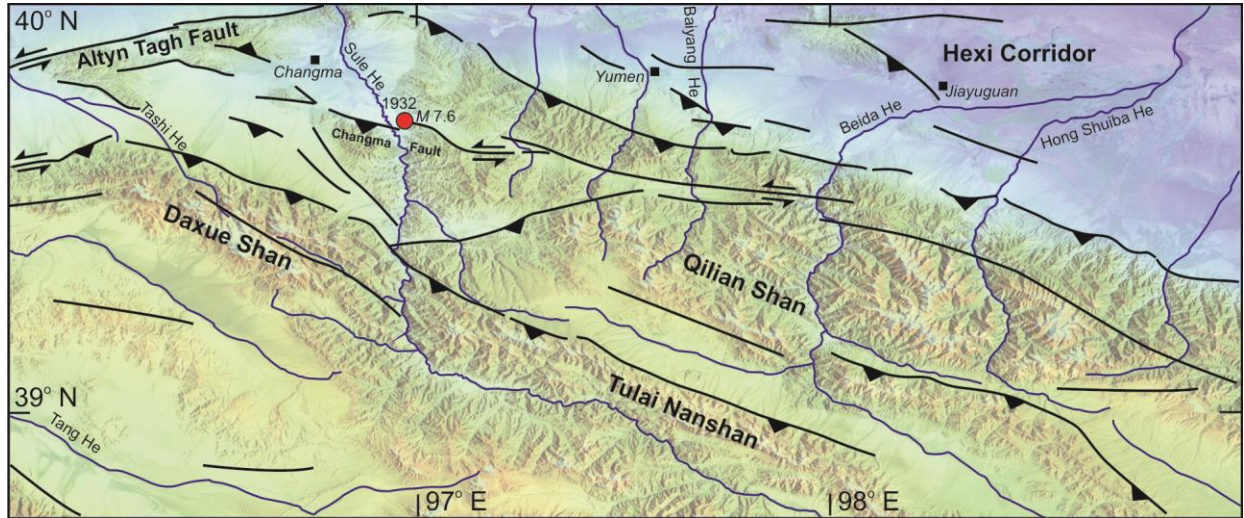


Figure 13. Drainage patterns and late Cenozoic faults in the northwest Qilian Shan. Background is SRTM digital topography superimposed over Landsat 7 false colour imagery (bands 7, 4, 2).

4 Discussion

4.1. Finite Strain estimates

Throw estimates of ~5 km for individual range fronts within the Qilian Shan can be summarised to estimate the overall range-normal shortening (Figure 5). We use an estimate of thrust dip of 45° , based on focal mechanism fault dips (Figure 4b) and the dips of exposed range-bounding thrusts (Figure 7). Thrusts are assumed to be planar, at least at the upper crustal levels under consideration. The throw corresponds to a horizontal shortening (heave) of ~5 km. Given the seven main internal range fronts in our longer section, this configuration suggests roughly 35 km of Cenozoic shortening across the section, or ~30%. Extrapolating across the full width of the Qilian Shan ranges (presently ~250-300 km, increasing westwards) gives ~110-130 km of thrust-related shortening in the Cenozoic within the interior of the Qilian Shan. An estimated 10-15% shortening is expressed in the folded strata of the intermontane basins (Section 3.1.1), equivalent to a further ~15 km shortening across the whole range.

The above estimate does not include the northern margin of the Qilian Shan, which differs from the internal range fronts in size of the topographic step down to the Hexi Corridor basins to the north (Figures 4a and 13), and the lower fault dip. A recent interpretation of seismic reflection profiles across this northern margin suggests typically ~30 km of shortening on this thrust front alone [Zuza *et al.*, 2016], on a thrust system dipping south at $\sim 15^\circ$, which is consistent with the 17° dip for the shear zone underlying the range derived from the elastic dislocation model in this study (Figure 11). Adding this 30 km value to the shortening estimate for the internal ranges gives a total of 155-175 km range-normal shortening across the Qilian

Shan. Previous interpretations of the amount of shortening at the northern, marginal, range front are lower than 30 km, at ≤ 10 km [Zheng *et al.*, 2010]. If this lower figure is correct, the overall shortening estimate must be reduced accordingly.

It has been previously suggested that at ~ 15 Ma there was the onset or acceleration of significant deformation and uplift in the northeastern part of the Tibetan Plateau, including the Qilian Shan [e.g. Duvall *et al.*, 2013]. Zheng *et al.* [2010] put the onset of significant late Cenozoic exhumation near the northern thrust in Qilian Shan to be slightly later, at 10 Ma, based on (U-Th)/He thermochronometry. Current GPS-derived deformation rates for convergence across the Qilian Shan are 7.7 mm yr^{-1} , combining dip-slip and strike-slip components (section 3.2.2). Extrapolating this rate to 15 Ma would mean ~ 115 km of total shortening across the Qilian Shan and Qaidam Basin, roughly split into ~ 65 km of strike-slip offset on the Haiyuan Fault, and ~ 100 km of range-normal shortening. The extrapolated range-normal shortening estimate is comparable to, if lower than, our estimate in this study of 155-175 km, and the estimate of 150-200 km shortening of Meyer *et al.* [1998]. It is distinctly lower than the >215 km derived by Zuza *et al.*, [2016] and the 250 km proposed by Cheng *et al.*, [2015]. The differences relate to the amount of strain interpreted for the internal ranges, which should clearly be a focus for future study. If the higher estimates of Zuza *et al.*, [2016] and Cheng *et al.*, [2015] are more realistic, strain needs to have taken place at near present rates for much longer than 15 Myr, or episodically through the Cenozoic. Our 155-175 km estimate fits a scenario where the present shortening rate has occurred since ~ 15 Ma, and achieved most of the finite strain, but leaves open the possibility for significant strain before 15 Ma.

4.2. Crustal scale accommodation of oblique convergence

Elastic dislocation modeling (Figure 11) provides an approach for understanding the deep structure of the region. The best-fit model for the interseismic geodetic data is consistent with localized, oblique, aseismic slip on a low-angle detachment thrust (17° dip) below a locking depth fixed of 26 km (Figure 14). A significant result of this model is that the locked-creeping transition lies directly beneath the location of the strike-slip Haiyuan Fault. This implies that strain is spatially partitioned in the seismogenic upper crust between strike-slip motion on the vertical Haiyuan Fault and thrusting, possibly to a large extent on the up-dip continuation of the detachment (Figure 11). This deformation pattern is similar to geodetic observations made across oblique oceanic subduction zones [McCaffrey *et al.*, 2000], and suggests that such behavior also occurs during strain partitioning in continental fold-and-thrust belts. It matches the model predictions of Bowman *et al* [2003]. A similar, but not identical, pattern of strain partitioning has been interpreted for the Haiyuan Fault and adjacent thrusts, to the east of our study area (102° - 104° E), by Daout *et al* [2016], based on a combined study of InSAR and GPS data. The two areas are different, in terms of the proportions of thrust and strike-slip deformation, and the greater depth of locking further west. A more fundamental difference is that Daout *et al* [2016] do not model the seismogenic strike-slip fault as directly above locked-creeping transition, as is found in our paper.

The historical events of $M \sim 7$ (Figure 4a) may be a record of rare but large earthquakes related to such a detachment, above the locking depth. It is not clear from the structures exposed within the Qilian Shan or our analysis of the geodetic data that the deep structure corresponds to a crustal scale critical wedge [Meyer *et al.*, 1998] or a continental subduction zone [Guillot and Replumaz, 2013]. The regional topography (Figures 4a, 11 and 12) shows steep margins on both

the northern and southern sides of the range, with lower internal slopes, not consistent with a crustal wedge thickening southwards towards the main part of the Tibetan Plateau.

It is notable that there is a clear instrumental seismicity record of relatively steep thrusts ($\geq 30^\circ$; Figure 4a), but no focal mechanisms that clearly correspond to seismic slip on the underlying detachment thrust. It could be that seismic slip on these steeper thrusts reduces the overall slip deficit on the detachment thrust, and reduces the magnitude of an eventual earthquake on the detachment.

We speculate that the detachment thrust corresponds to the upper surface of the North China Block (Figure 3), proposed by some authors [e.g. *Yin et al.*, 2008b] to underlie most or all of the Qilian Shan, and consistent with deep seismic evidence [*Ye et al.*, 2015]. Given the lack of hundred km scale Cenozoic thrusting at the northern margin of the Qilian Shan [*Hetzl et al.*, 2004; *Zheng et al.*, 2010; *Zuza et al.*, 2016; this study], we conclude that the main underthrusting of the North China Block (Figure 3) is more likely to have taken place during the Paleozoic accretion of the Qilian Shan, with reactivation during late Cenozoic deformation. There is evidence for this Paleozoic age for underthrusting, in the zircon populations within Paleozoic plutons in the Qilian Shan, which contain Precambrian signatures of the North China Block [*Gehrels et al.*, 2003].

The structure of the Qilian Shan has similarities to the Himalayas, where the Indian Plate underthrusts Eurasian crust along a major thrust, dipping at a similar angle [e.g. *Murphy et al.*, 2014]. There is no evidence that the underthrust North China Block underwent slab break-off and uplift in the Cenozoic, comparable to the India-Eurasia system [*Magni et al.*, 2017].

4.3. Thrust distribution, crustal thickening and surface uplift

There is a concentration of $M > 5$ thrust events at or below the regional 3500 m elevation contour (Figure 4a). Of the few $M > 5$ thrust events that occurred at higher elevations, one occurs on a restraining bend along the Haiyuan Fault, and two are ~ 100 km from the Altyn Tagh Fault. $M > 4$ events occur at higher elevations (Figure 4a), but mainly on the northern side of the range. The area west of Lake Hala is distinctly aseismic. Whilst no focal mechanisms are available for these smaller events, it seems reasonable to conclude that they are mainly thrust earthquakes, in keeping with the observed structure of the region. Events along the Haiyuan Fault are more likely to be strike-slip. The higher, interior regions of the Qilian Shan show less evidence of Holocene thrust activity than lower elevation regions (Figure 4a). The interior of the Qilian Shan has more subdued relief than the range margins [*Liu-zeng et al.*, 2008] (Figure 12). The combined low relief and low seismicity give the interior region plateau-like characteristics, at least in incipient form.

4.4 Influence of the Altyn Tagh Fault

It has been previously proposed that thrusts in the Qilian Shan are splay off the northeastern part of the Altyn Tagh Fault, for up to ~ 400 km away from the latter, and owe their existence to the need to absorb the slip along this structure. [*Meyer et al.*, 1998; *Van der Woerd et al.*, 2001; *Cheng et al.*, 2015]. There is clearly interaction between the thrusts in the Qilian Shan and the Altyn Tagh Fault: the thrusts in the northwest of the Qilian Shan change strike westwards, to swing into alignment (Figure 2). Active thrusting occurs relatively close to the

669 fault, at high regional elevations, in contrast to the remainder of the Qilian Shan further east
670 (Figure 4a).

671 However, the regional elevations of the Qilian Shan decrease towards the Altyn Tagh
672 Fault (Figures 4a and 12), implying that the greatest crustal thickening and highest topography
673 has developed in the interior of the Qilian Shan, rather than dying away laterally from the Altyn
674 Tagh Fault. The pattern of rivers in the northwestern Qilian Shan is consistent with an origin
675 whereby drainage flows westwards away from the interior of the range, before meeting uplifted
676 areas influenced by the Altyn Tagh Fault, and being diverted northwards towards the Hexi
677 Corridor (Figure 13). Therefore the influence of the Altyn Tagh Fault may extend laterally only
678 100-200 km into the interior of the Qilian Shan, with most of the thrusting along the range
679 representing a regional shortening as part of the India-Eurasia convergence. There is clearly a lot
680 of scope for more work on this subject, but our observations in Figures 12 and 13 are consistent
681 with the other results in this paper, whereby the thrusting in the Qilian Shan is a major part of the
682 overall India-Eurasia convergence, and not a secondary feature arising from the slip along the
683 Altyn Tagh Fault.

684 We also note the coincidence of Cenozoic thrusts with Paleozoic structural boundaries
685 along the length of the Qilian Shan [Song *et al.*, 2013; Wu *et al.*, 2016], i.e. for ~800 km, and
686 strongly implicating reactivation of basement faults as a factor in controlling the location of the
687 young structures (Figure 2). Our fieldwork observations confirm this match, with the exception
688 of the Xitieshan region (Figure 7), where the Paleozoic foliations strike more north-south than
689 the Cenozoic thrusts in the same region.

691 5 Conclusions

692 We estimate range-normal shortening across the Qilian Shan as ~155-175 km,
693 comparable to the 150-200 km claimed from crustal scale restorations by Meyer *et al.*, 1998], but
694 lower than the >215 km and 250 km values derived by Zuza *et al.*, [2016] and Cheng *et al.*
695 [2015]. Maintaining the present range-normal shortening rate of 6.4 mm yr⁻¹ for the 15 Myr since
696 the inferred onset or acceleration of deformation in the Qilian Shan yields ~100 km total range-
697 normal shortening, which is to a first order consistent with our estimate, but raises the possibility
698 that some strain took place before 15 Ma, or has even slowed since 15 Ma.

699 The overall structure of the Qilian Shan is divergent, with southward-thrusting at the
700 southern margin and northward-thrusting at the northern margin of the range. Thrust structure
701 within the region shows no predominant transport direction at upper crustal levels (Figure 5).

702 Elastic dislocation modeling indicates that a single regional detachment thrust (dip angle
703 of 17°) with oblique slip below 26 km depth can explain both range-parallel and range-normal
704 components of the GPS-derived interseismic velocity field (Figure 11). Upper and lower crustal
705 deformation therefore appear to be decoupled along a zone approximating simple shear under the
706 Qilian Shan (Figure 14).

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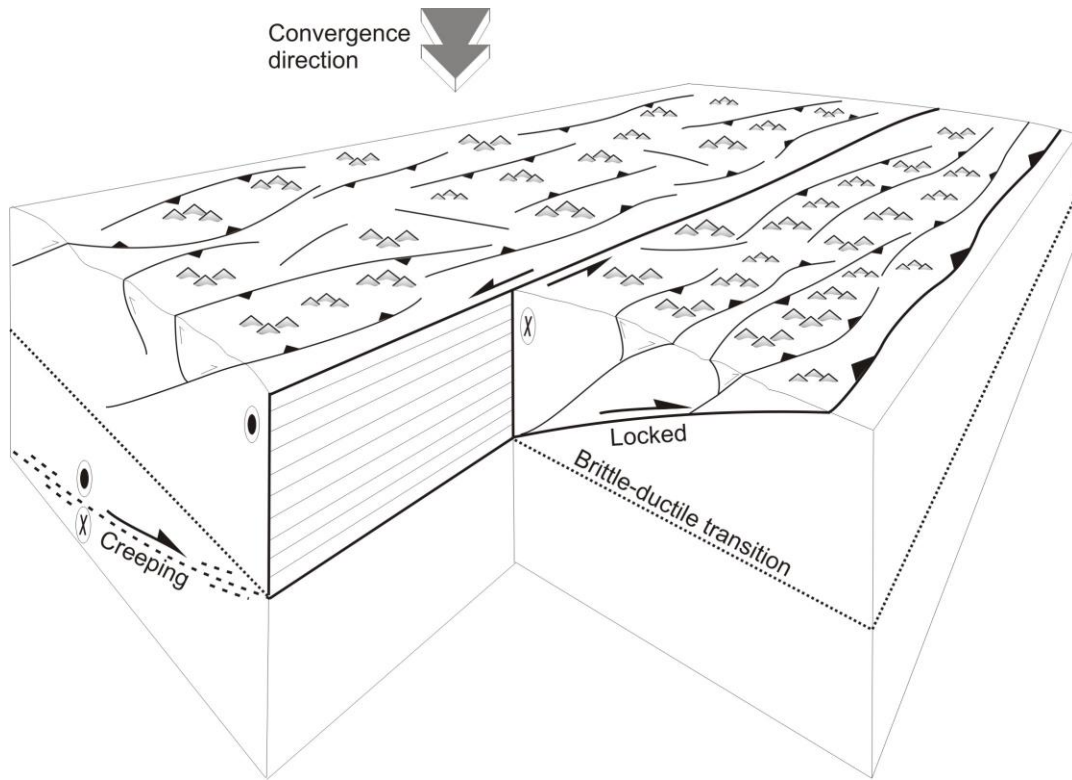


Figure 14. Schematic illustration of the fault geometry proposed for the Qilian Shan, with partitioning of oblique convergence onto strike-slip and thrust components, with the intersection of the strike-slip and thrust faulting taking place at the down-dip limit of the locked portion of the thrust.

This shear zone may be the upper surface of an underthrust North China Block [Yin *et al.*, 2008b] (Figure 3), with the caveat that regional underthrusting is more likely to have taken place in the Paleozoic than the Cenozoic. The dip angle is higher than commonly proposed critical wedge scenarios for individual fold-and-thrust belts, which have much lower dip values for the basal detachment to the fold-and-thrust belt [Davis *et al.*, 1983].

The strike-slip component of deformation is located on a near-vertical structure above the up-dip limit of oblique interseismic slip on the model detachment, and equivalent to the observed Haiyuan Fault. This configuration of strain partitioning may apply more generally in zones of oblique convergence (Figure 14), and so explain the location of strike-slip faults in other fold-and-thrust belts, both ancient and active [e.g. Holdsworth and Strachan, 1991].

Well-constrained thrust focal mechanisms dip typically at $>30^\circ$ (Figure 4b) and are confined to the upper 20 km of the crust [Sloan *et al.*, 2011]. If the crustal-scale structure involves an underlying low-angle thrust (Figures 3 and 11), it has not produced earthquakes recorded in the instrumental catalogs. However, historic $M \sim 7$ events at the northern side of the Qilian Shan could correspond to seismic slip on this structure (Figure 4a).

Major thrust seismicity ($M > 5$) in the Qilian Shan diminishes above the regional 3500 m contour, and the few large thrust events above this elevation are typically linked to strike-slip faults at restraining bends (Figure 4a). Smaller earthquakes are more widespread across the

Qilian Shan (Figure 4a). Combined with the subdued relief at these higher elevations (Figure 12), it appears that the interior of the Qilian Shan is developing plateau-like characteristics.

The influence of the Altyn Tagh Fault may be confined to adjacent parts of the Qilian Shan, within 100-200 km, and not the full lateral extent of thrust ranges. These Cenozoic thrusts follow the Paleozoic basement grain for ~800 km to the ESE, along the Qilian Shan [Song *et al.*, 2013; Wu *et al.*, 2016], and strongly imply that reactivation of Paleozoic structures controls the location of the modern faults.

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